Development of a Simplified Adaptive Finite Element Model of the Gulf Stream

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Abstract

The goal of this dissertation is to look at the feasibility of incorporating a moving adaptive mesh refinement procedure into an existing finite element model of the shallow water equations. The Stommel model simulates simple oceanic flows including western boundary currents like the Gulf Stream. This dissertation describes the finite element formulation of the Stommel equations, a modified form of the shallow water equations. The feasibility of including a mesh movement procedure based on monitor functions is looked at, with a study of one particular monitor. Results are shown for a specific mesh velocity and for mesh velocities computed from this monitor. We conclude that the monitor function used is a promising start at finding a way to allow the mesh to dynamically evolve with the flow, in applications such as this.

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Declaration

I confirm that this work is my own and the use of all other material from other sources has been properly and fully acknowledged.

Nicola Stone.

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Chapter 1

Introduction

Traditionally, ocean models used structured grids, reflecting the computational resources that were available at the time. Now ocean modellers are starting to use unstructured and adapting meshes more.

Unstructured meshes have many advantages over structured meshes. The coastline and basin geometry can be better represented. Also, the mesh can be easily adapted to give an increased local mesh resolution, or a dynamically adaptive mesh [4]. In the first case, a high resolution grid can be placed in an area of interest with lower resolution in less important areas, to give optimal use of computer power. In a North Atlantic model, for example, we might place a higher resolution grid in the region of the Gulf Stream. In the second case, we can have a mesh which moves to follow certain aspects of the solution, for example eddies moving across a basin.

Adaptive meshes are useful as we may not know the positions of developing flow features before the model is run. The resolution of the model can be changed to optimally resolve the flow. Adaptive methods can also be incorporated into existing models and may give increased 2 Introduction

efficiency and accuracy of the solution.

This dissertation will investigate the feasibility of using adaptive mesh movement in the Stommel model. The mesh movement is determined by following a certain feature (fluid volume) of the solution in Stommel's equations of ocean circulation with a prescribed vorticity. We will be looking at the possibility of including mesh movement into an existing two-dimensional finite element model. The existing model uses an implicit time discretisation with a black box solver the solve the equations. To avoid having to invert a matrix, this has been changed to an explicit scheme for this dissertation. The mesh movement is based on the idea of having a monitor function to determine how the mesh will move. The monitor function we have chosen is the depth of the ocean $H = h + \eta$, where h is the resting depth of the fluid and η is the surface elevation, and the prescribed vorticity of the mesh is taken to be zero.

We start by stating the shallow water equations in Chapter 2. Then we briefly describe how Stommel simplified these equations to create a model which describes the creation of western boundary layers in areas of the ocean such as the Gulf Stream. Scale analysis is performed on these equations to determine the dominant terms, which allow us to make assumptions of the nature of the solution without first having to solve the equation.

In Chapter 3 we look at how the governing equations of Stommel's model are formulated using finite elements, and then discretised in time [1]. We then describe how the solution is computed. The initial conditions of the model are stated, along with the initial mesh. The elevation solution on a fixed mesh is then described as a motivation to

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explore the possibility of moving the mesh to follow a flow.

The idea of generating a mesh velocity from the conservation of a monitor function [8] is introduced in Chapter 4. We have chosen one particular monitor function which ensures that the volume of fluid in each element remains constant over all time. This is one of many monitor functions that could be used to move the mesh. Likewise, the choice of zero mesh vorticity is only one of many choices that can be made.

Then an equation for the mesh velocity induced by this monitor function is formulated using finite element methods in Chapter 5. The entries of the resulting stiffness matrix and load vector are then derived, and a mesh velocity calculated for each element.

Chapter 6 describes how this element velocity is interpolated onto the nodes, and the mid-segments of the elements in the mesh. The mesh node velocity is then used in a mesh movement equation to move the mesh. As the true solution should not change when the mesh is modified, the mid-segment velocity is used in advection terms included in the governing equations to compensate for the mesh movement.

The results of numerical simulations carried out to move the mesh are shown in Chapter 7. These test situations look at the suitability of different ideas for mesh movement. The first test carried out is a simple prescribed westerly mesh movement, given that we know the solution travels to the west. Then the results for the conservation of volume principle are discussed. The next test involves the Lagrangian idea of moving the mesh according to the flow velocity which is a special case of the computed velocity when the vorticity of the mesh is equal to the vorticity of the fluid. The last test is a linear combination of

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the monitor function mesh velocity and the flow velocity based mesh movement.

Chapter 8 draws conclusions on the methods used within the dissertation, and on the feasibility of using the conservation of a monitor function to prescribe the mesh movement. Problems with the tests in chapter 7 are also addressed.

Finally in Chapter 9, ideas for further work are discussed. These are ideas that would have been interesting to investigate, had the dissertation not been limited by time constraints.

Chapter 2

The Stommel Model

This chapter will describe the Stommel model, looking at the governing equations and how the model is represented using finite element methods. Then the time discretisation will be described, along with how the results are computed. This is followed by results of a fixed mesh finite element model of the Stommel model in a square domain.

2.1 The Shallow Water Equations

The shallow water equations on domain Ω with boundary $d\Omega$

$$\frac{\partial \eta}{\partial t} + \nabla \cdot (H\mathbf{u}) = 0, \tag{2.1}$$

$$\frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} + f \mathbf{e}_z \times \mathbf{u} = -g \nabla \eta + \frac{\nu}{H} \nabla \cdot (H \nabla \mathbf{u}) + \frac{\tau^{\eta} - \tau^b}{\rho_0 H}, \tag{2.2}$$

are solved for $\eta(\mathbf{x}, t)$, the surface elevation, and $\mathbf{u}(\mathbf{x}, t)$, the depth-averaged velocity. In these equations $H = h + \eta$ is the depth of the fluid, where $h(\mathbf{x})$ is the resting depth of the fluid, as shown in Figure 2.1, \mathbf{e}_z is a unit vector in the vertical direction, g is the gravitational

acceleration, ν is the eddy viscosity, τ^{η} and τ^{b} are the surface and bottom stresses respectively and ρ_{0} is the homogeneous density of the fluid.

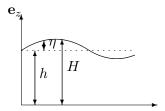


Figure 2.1: Depth of the fluid (1-D)

Equations (2.1) and (2.2) describe the development of an incompressible flow of constant density on the surface of the sphere. They are only relevant when the scale of the region of flow is much larger than the depth of the fluid, as in the oceans where the depth is much smaller in proportion to the size of the ocean. Equation (2.1) describes conservation of mass over the domain, and (2.2) describes the momentum balance.

2.2 The Stommel Model

The Stommel model is a simplification of these shallow water equations. It contains two important forces acting on ocean flow, rotational forces and wind effects, and allows us to model the wind-driven ocean circulation. An important effect of wind driven circulation in the ocean basins is the strong western boundary currents. The Gulf Stream in the North Atlantic is one such current, transporting warm tropical water to higher latitudes.

The shallow water equations of the previous section describe the motion of fluid on the surface of a sphere. The Stommel model simplifies this by considering the motion of fluid in Cartesian coordinates on an f-plane, where to first order the Coriolis parameter is constant.

Stommel [5] tried to understand the cause of the intensification of these boundary currents, by proposing the following

- considering the Coriolis parameter changing with latitude at a constant rate by letting $f = f_0 + \beta y$ (β -plane approximation.)
- including a bottom friction term, so keeping the higher order terms and the ability to satisfy conditions of no normal flow over the boundaries.

Then, also neglecting non-linear and diffusion terms, the governing equations of the Stommel model become

$$\frac{\partial \eta}{\partial t} + h \nabla \cdot \mathbf{u} = 0, \tag{2.3}$$

$$\frac{\partial \mathbf{u}}{\partial t} + (f_0 + \beta y)\mathbf{e}_z \times \mathbf{u} = -g\nabla \eta - \gamma \mathbf{u} + \frac{\tau^{\eta}}{\rho h}.$$
 (2.4)

In these equations, $\mathbf{x} = (x, y)$ is the spatial coordinate, h is the resting depth of the fluid which is assumed constant, f_0 is the reference value of the Coriolis parameter first derivative in the y direction, \mathbf{e}_z is a unit vector in the vertical direction, g is the gravitational acceleration, γ is a linear friction coefficient, ρ is the homogeneous density of the fluid, τ^{η} is the wind stress acting on the surface of the fluid and ∇ is the two dimensional gradient operator. The domain Ω is the interval $[0,L] \times [0,L]$. Stommel assumed a simple form for the wind stress

$$\tau_x = -\tau_0 \cos\left(\frac{\pi y}{y_n}\right), \quad 0 \le y \le L,$$

$$\tau_y = 0.$$

These equations are solved subject to no-normal flow boundary conditions, $\mathbf{u} \cdot \mathbf{n} = 0$ on $\partial \Omega$.

We note that the there is an analytic solution to the Stommel equation, so we know that the solution moves westward to create a boundary layer at the western boundary. We might then think that we could just prescribe the mesh movement to be to the west for all time. The results for this idea are shown in Chapter 7.

2.3 Scale Analysis

We now use scale analysis to compare the important terms in the Stommel equations in relation to the other terms. This gives us an insight into the flow of the fluid before we solve the equations. It may also help us to determine possibilities for monitor functions that could be used to move the mesh. The parameters we will use for scaling are shown in Table 2.1.

The existing code that we are using solves the dimensionless shallow water equations. The dimensional variables are multiplied by a scale factor to make them dimensionless. If we choose scale factors carefully, we end up with a set of equations where all variables and coordinates are of order 1. Scaling is as follows:

$$x' = L_h x,$$

$$u' = Uu,$$

$$h' = L_v h,$$

Parameter	Description	Scale
E	vertical surface elevation	1 m
U	horizontal velocity	$0.1 \; \mathrm{m \; s^{-1}}$
L_h	horizontal length	$1 \times 10^6 \text{ m}$
L_v	vertical length	$1 \times 10^3 \text{ m}$
T	time scale	$1 \times 10^4 \mathrm{\ s}$
f	Coriolis term	$1 \times 10^{-4} \text{ s}^{-1}$
β	beta term	$1 \times 10^{-11} \text{ m}^{-1} \text{ s}^{-1}$
g	gravitational term	$10 \text{ m}^2 \text{ s}^{-1}$
γ	drag coefficient	1×10^{-4} (dimensionless)
ρ	density	$1\times10^3~\rm kg~m^{-3}$
au	wind stress	$0.2 \text{ kg m}^{-1} \text{ s}^{-2}$

Table 2.1: Scales for each parameter in the Stommel equations

$$\eta' = E, \eta$$

$$t' = Tt,$$

where ' denotes a dimensional variable. These dimensional variables are then replaced in the equations by their dimensionless expressions.

The continuity equation then reads

$$\frac{E}{T}\frac{\partial \eta}{\partial t} + \frac{L_v U}{L_b} h \nabla \cdot \mathbf{u} = 0.$$

We can then evaluate the relative importance of each term in the equation by dividing each term by the same dimensional factor. For this equation, we divide through by $\frac{E}{T}$ to obtain

$$\frac{\partial \eta}{\partial t} + \frac{L_v U T}{L_h E} h \nabla \cdot \mathbf{u} = 0,$$

where $\frac{L_vUT}{L_hE}$ is a dimensionless number. If this was large then the divergence term would be of more importance than the time derivative.

We want these two terms to be of equal importance as it would not be sensible to neglect one of the two terms in the equation. Therefore in the code we choose $T = \frac{L_h E}{L_v U}$, so that both terms are of equal importance in the equation (as the dimensionless number is equal to 1). This is a way of determining a timescale, which we do not know exactly beforehand.

We do the same for the momentum equation to obtain

$$\frac{U}{T}\frac{\partial \mathbf{u}}{\partial t} + U\left(f_0 + \beta L_h y\right) \mathbf{e}_z \times \mathbf{u} = -\frac{E}{L_h} g \nabla \eta - U \gamma \mathbf{u} + \frac{\tau}{\rho L_n} \frac{\tau^{\eta}}{\rho h}.$$

This time we divide through by the dimensional factor in front of the Coriolis term (f_0U) . Therefore we obtain the dimensionless equation:

$$\frac{1}{f_0 T} \frac{\partial \mathbf{u}}{\partial t} + \left(1 + \frac{\beta L_h}{f_0} y\right) \mathbf{e}_z \times \mathbf{u} = -\frac{gE}{f_0 U L_h} \nabla \eta - \frac{\gamma}{f_0} \mathbf{u} + \frac{\tau}{\rho f_0 U L_v} \frac{\tau^{\eta}}{\rho h}.$$

In the existing code, U is chosen so that $\frac{gE}{f_0UL_h} = 1$. This means that $U = \frac{gE}{f_0L_h}$. The remaining scale factors are shown in Table 2.1. With these scales, we can see that the dimensionless numbers in front of the time derivative term, the gravitational term and the drag term are all equal to 1. The leading terms in this equation are the Coriolis term (dimensionless number equal to 1.1) and the pressure gradient term (dimensionless number equal to 2). This suggests that the rotational property of the flow velocity and wind stresses are dominant features of these equations. We might then consider using the flow velocities as a basis for mesh movement. Results for using this Lagrangian idea are presented and discussed in Chapter 7.

Chapter 3

Numerical Solution to the Stommel Problem

We now look at, following Hanert [1], how the Stommel problem is formulated using the finite element method.

3.1 Finite Element Formulation

First we multiply each equation by a test function and then integrate over the domain. The test functions we use are $\hat{\eta}$ for the continuity equation and $\hat{\mathbf{u}}$ for the momentum equation. This gives us the weak formulation of equations (2.3) and (2.4) on Ω as

$$\int_{\Omega} \frac{\partial \eta}{\partial t} \hat{\eta} d\Omega + h \int_{\partial\Omega} \mathbf{u} \cdot \mathbf{n} \hat{\eta} d\Gamma - h \int_{\Omega} \mathbf{u} \cdot \nabla \hat{\eta} d\Omega = 0,$$

$$\int_{\Omega} \frac{\partial \mathbf{u}}{\partial t} \hat{\mathbf{u}} d\Omega + \int_{\Omega} (f + \beta y) (\mathbf{e}_{z} \times \mathbf{u}) \hat{\mathbf{u}} d\Omega$$

$$+g \int_{\Omega} \nabla \eta \hat{\mathbf{u}} d\Omega + \gamma \int_{\Omega} \mathbf{u} \hat{\mathbf{u}} d\Omega - \int_{\Omega} \frac{\tau^{\eta}}{\rho h} \hat{\mathbf{u}} d\Omega. = 0,$$
(3.1)

where the divergence term has been integrated by parts and the resulting boundary integral can therefore be removed as $\mathbf{u} \cdot \mathbf{n} = 0$ on $\partial \Omega$.

To make the finite element approximation, we use a mixture of conforming (P_1) and nonconforming basis functions (P_1^{NC}) . The conforming nodes are on the vertices of the triangulation, whilst the nonconforming nodes are on the mid segment of the triangulation (Figure 3.1). This P_1^{NC} - P_1 finite element pair is very good for solving the shallow water equations [3]. This is due to the orthogonality of the shape functions, which allows us to reduce the computational cost of the scheme. The nonconforming elements are discontinuous everywhere, except at the mid segments. They are a compromise between continuous and discontinuous approximations.

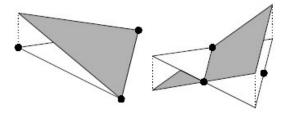


Figure 3.1: P_1 and P_1^{NC} basis functions

We replace \mathbf{u} and η in (3.1) and (3.2) by their discrete approximations \mathbf{u}^h and η^h as follows. Let

$$\eta pprox \eta^h = \sum_{i=1}^{N_V} \eta_i \phi_i,$$

 $\mathbf{u} pprox \mathbf{u}^h = \sum_{j=1}^{N_S} \mathbf{u}_j \psi_j,$

where η_i and \mathbf{u}_j represent the elevation and velocity nodal values, ϕ_i and ψ_j represent the elevation and velocity basis functions, and N_V

and N_S are the number of vertices and segments in the mesh. The nodal values are then calculated using the Galerkin method; replacing $\hat{\eta}$ by ϕ_i ($1 \le i \le N_V$), and $\hat{\mathbf{u}}$ by $(\psi_j, 0)$ and $(0, \psi_j)$ ($1 \le j \le N_S$) in (3.1) and (3.2). This results in

$$\frac{d\eta_{j}}{dt} \int_{\Omega} \phi_{i}\phi_{j}d\Omega - h\mathbf{u}_{j} \int_{\Omega} \nabla\phi_{i}\psi_{j}d\Omega = 0, \quad (3.3)$$

$$\frac{d\mathbf{u}_{j}}{dt} \int_{\Omega} \psi_{i}\psi_{j}d\Omega + (f_{0} + \beta y_{j})(\mathbf{e}_{z} \times \mathbf{u}_{j}) \int_{\Omega} \psi_{i}\psi_{j}d\Omega$$

$$+g\eta_{j} \int_{\Omega} \psi_{i}\nabla\phi_{j}d\Omega + \gamma\mathbf{u}_{j} \int_{\Omega} \psi_{i}\psi_{j}d\Omega - \int_{\Omega} \frac{\tau^{\eta}}{\rho h}\psi_{i}d\Omega = 0, \quad (3.4)$$

Note that in the second term of equation (3.4) we have also approximated $y \approx y_i$.

3.2 Temporal Discretisation

We need to also discretise equations (3.3) and (3.4) in time. For a given time step $\Delta t = t^{n+1} - t^n$, we have the following equations.

$$\frac{\eta^{n+1} - \eta^n}{\Delta t} \int_{\Omega} \phi_i \phi_j d\Omega = h \mathbf{u}_j \int_{\Omega} \nabla \phi_i \psi_j d\Omega$$

$$\equiv f_{\eta}(\eta, \mathbf{u}), \qquad (3.5)$$

$$\frac{\mathbf{u}^{n+1} - \mathbf{u}^n}{\Delta t} \int_{\Omega} \psi_i \psi_j d\Omega = -(f_0 + \beta y_j)(\mathbf{e}_z \times \mathbf{u}_j) \int_{\Omega} \psi_i \psi_j d\Omega$$

$$-g \eta_j \int_{\Omega} \psi_i \nabla \phi_j d\Omega - \gamma \mathbf{u}_j \int_{\Omega} \psi_i \psi_j d\Omega$$

$$+ \int_{\Omega} \frac{\tau^{\eta}}{\rho h} \psi_i d\Omega$$

$$\equiv f_{\mathbf{u}}(\eta, \mathbf{u}), \qquad (3.6)$$

Where f_{η} and $f_{\mathbf{u}}$ are used in the Adams Bashforth scheme as described below.

We now use the third order Adams Bashforth scheme to solve these equations for η and \mathbf{u} . This gives us two matrix systems

$$A_{\eta}\eta^{n+1} = \mathbf{b}_{\eta},$$
$$A_{\mathbf{u}}\mathbf{u}^{n+1} = \mathbf{b}_{\mathbf{u}},$$

where

$$A_{\eta}^{i,j} = \int_{\Omega} \phi_i \phi_j d\Omega,$$

$$A_{\mathbf{u}}^{i,j} = \int_{\Omega} \psi_i \psi_j d\Omega,$$

and

$$\mathbf{b}_{\eta}^{i} = \sum_{j} \left(\eta_{j}^{n} \int_{\Omega} \phi_{i} \phi_{j} d\Omega + \frac{\Delta t}{12} \left(23 f_{\eta}^{n} - 16 f_{\eta}^{n-1} + 5 f_{\eta}^{n-2} \right) \right),$$

$$\mathbf{b}_{\mathbf{u}}^{i} = \sum_{j} \left(\mathbf{u}_{j}^{n} \int_{\Omega} \psi_{i} \psi_{j} d\Omega + \frac{\Delta t}{12} \left(23 f_{\mathbf{u}}^{n} - 16 f_{\mathbf{u}}^{n-1} + 5 f_{\mathbf{u}}^{n-2} \right) \right).$$

These can be solved by inverting the mass matrices A_{η} and $A_{\mathbf{u}}$.

3.3 Computation

We now discuss how the entries of the mass matrices are computed to simplify the calculation of the solution.

3.3.1 Mass Matrix Entries

The entries of the mass matrices A_{η} and $A_{\mathbf{u}}$ for each element are

$$\begin{array}{lcl} A_{\eta}^{i,j} & = & \int\limits_{\Omega_e} \phi_i \phi_j d\Omega, \\ \\ A_{\mathbf{u}}^{i,j} & = & \int\limits_{\Omega} \psi_i \psi_j d\Omega. \end{array}$$

So we have for each element

$$A_{\eta}^{i,j} = \frac{\Delta_{\text{area}}}{12} \begin{pmatrix} 2 & 1 & 1 \\ 1 & 2 & 1 \\ 1 & 1 & 2 \end{pmatrix}, \qquad (3.7)$$

$$A_{\mathbf{u}}^{i,j} = \frac{\Delta_{\text{area}}}{3} \begin{pmatrix} 1 & 0 & 0 \\ 0 & 1 & 0 \\ 0 & 0 & 1 \end{pmatrix}.$$

These are then assembled to form mass matrices for the whole domain.

An alternative to inverting these large matrices is to note that $A_{\mathbf{u}}$ is a diagonal matrix, so when computing the solution we can think of it as a vector $\mathbf{A}_{\mathbf{u}}^{diag}$ containing all of the diagonal entries of $A_{\mathbf{u}}$. A_{η} however is not diagonal so it is not so simple to create a vector in this way. However, if we assume the mass of each element is concentrated at its nodes then the matrix can be converted into a diagonal matrix by lumping the off diagonal entries in A_{η} onto the diagonal. Then this diagonal matrix can be represented by the vector \mathbf{A}_{η}^{diag} .

3.3.2 Computing the solution

To compute the solution it is now easier to represent the two systems as one system

$$\mathbf{A}^{diag} \left(egin{array}{c} dots \ \eta_j \ dots \ \mathbf{u}_j \ dots \end{array}
ight)^{n+1} = \mathbf{b},$$

where \mathbf{A}^{diag} contains all of the entries from \mathbf{A}^{diag}_{η} and $\mathbf{A}^{diag}_{\mathbf{u}}$, and \mathbf{b} contains all of the entries from \mathbf{b}_{η} and $\mathbf{b}_{\mathbf{u}}$. We can then compute the

solution for η_j and \mathbf{u}_j simply by dividing \mathbf{b} by \mathbf{A}^{diag} .

3.4 Initial Conditions

Initially the solution is zero everywhere in Ω for the elevation and velocity fields. The movement of the solution is determined by the wind stress function which is set to

$$\tau^{\eta} = \begin{pmatrix} 0.2\sin\left(y - \frac{L}{2}\right) \\ 0 \end{pmatrix}.$$

The initial mesh shown in Figure 3.2 is a structured mesh with equal sized elements

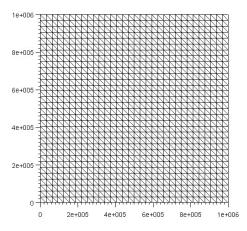


Figure 3.2: Initial Mesh

When using the third order Adams Bashforth scheme for the temporal discretisation of the Stommel equations, the time step that is required to keep the solution stable is $\Delta_t = 20$ seconds. We then run the model with this time step and let the model converge to a steady solution.

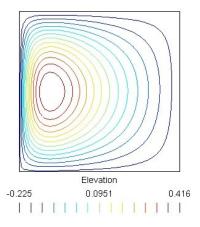


Figure 3.3: Stommel solution on a square domain

3.5 Stommel Solution

The results for the Stommel problem show a strong western intensification (Figure 3.3). The flow created by the wind stress is initially in the same direction as the wind stresses, from the west to the east in the top half of the region, and from the east to the west in the bottom half of the region. The flow starts to rotate around the centre of the region. Then because of the change in Coriolis parameter with latitude, the solution moves west creating the boundary current. The flow is southward for most of the region, and then returns in a swift northward jet on the western boundary. This motion idealises the Gulf Stream in which a narrow western boundary current returns the southward interior flow to the north. If Stommel had not included the Coriolis term in his equations, the solution would be centred in the middle of the domain with no westerly movement.

This solution leads us to explore a mesh that moves with the flow.

As the flow moves towards the west, we wish to move the mesh with it to increase the resolution in the developing boundary layer.

Chapter 4

Moving Mesh Adaptation

To adapt the mesh as the solution changes, we aim to find a velocity for each node in the mesh. This section will look at the principles behind the moving mesh, and how the mesh adaptation principle is formulated to calculate the mesh velocities.

4.1 Moving Mesh Methods

Moving mesh methods, or r-adaptive methods, move the positions of a fixed number of nodes in the mesh so that the nodes follow a feature or become concentrated in certain areas of the solution. In principle these methods are easy to implement and fixed mesh models can be easily adapted to incorporate moving adaptive meshes. As there is no addition and deletion of nodes, difficulties in restarting the time integration procedure are avoided. Another benefit is the ability of the mesh to move in a quasi-Lagrangian manner. If the mesh points were to move with the flow velocity, this would lead to reduced transport velocity in relation to the mesh, with the possibility of using large time

steps.

There are also drawbacks to these methods. The number of nodes is constant, which may not always be optimal because the level of dynamic behaviour may vary during the simulation. There is also the possibility that inappropriately shaped or tangled meshes may occur.

There are many ways of moving the mesh. We will be looking at a moving mesh method which creates mesh velocities based on monitor functions as described by Wells [8]. We require a monitor function to control the relative density of mesh points. The method we will discuss is based on a conservation principle, where the monitor function is to be conserved in time, which will then lead to mesh movement. A Eulerian conservation law can be derived from the conservation principle. We can obtain a unique mesh velocity when this conservation law is used along with a curl condition, which gives the rotational properties of the mesh.

4.2 Mesh Velocity Principle

We now following [7] with (h + eta) as the choice of monitor function.

In this project, the principle behind the mesh movement is to find a velocity $\dot{\mathbf{x}}$ such that the volume of water in each patch Ω_c of the domain Ω is constant in time. That is,

$$\int_{\Omega_c} (h + \eta) d\Omega = \text{constant in time}, \tag{4.1}$$

is true for each element. From this it follows that

$$\frac{d}{dt} \int_{\Omega_c} (h + \eta) d\Omega = 0. \tag{4.2}$$

Using the Reynolds Transport Theorem we can change this Lagrangian form of the conservation principle into an Eulerian form. This Theorem states that the rate of change of the integral of a property, N, over a moving control volume is equal to the integral of the rate of change of N over the control volume plus the net flux of N through the control surface. So we can rewrite (4.2) as

$$\int_{\Omega_c} \frac{\partial (h+\eta)}{\partial t} d\Omega + \oint_{\partial \Omega_c} (h+\eta) \dot{\mathbf{x}} \cdot \mathbf{n} d\Gamma = 0,$$

where $\dot{\mathbf{x}}$ is the boundary velocity. As h is constant it can be removed from the first integral. Also, as η is small compared with h, we replace $h + \eta$ by h in the second integral. Then using Gauss' Theorem on the second integral we obtain

$$\int_{\Omega_c} \left(\frac{\partial \eta}{\partial t} + \nabla \cdot (h\dot{\mathbf{x}}) \right) d\Omega = 0, \tag{4.3}$$

where $\dot{\mathbf{x}}$ is any velocity consistent with the boundary velocity. If we use the continuity equation (2.3) from the shallow water equations, we can replace $\frac{\partial \eta}{\partial t}$ in (4.3) to give

$$\int_{\Omega_c} \nabla \cdot (-h\mathbf{u} + h\dot{\mathbf{x}}) \, d\Omega = 0.$$

It follows that

$$\nabla \cdot (-h\mathbf{u} + h\dot{\mathbf{x}}) = 0.$$

and so

$$\nabla \cdot (h\dot{\mathbf{x}}) = \nabla \cdot (h\mathbf{u}).$$

This equation is not sufficient to determine $\dot{\mathbf{x}}$ but if we use Helmholtz decomposition Theorem, and let $\nabla \times \dot{\mathbf{x}} = 0$, we have that

$$\dot{\mathbf{x}} = \nabla p,\tag{4.4}$$

where p is the mesh velocity potential. We then have

$$\nabla \cdot (h\nabla p) = \nabla \cdot (h\mathbf{u}),\tag{4.5}$$

which has a unique solution, given p on the boundary or given $\frac{\partial p}{\partial n}$ on the boundary and one prescribed value of p.

Chapter 5

Solving for the Mesh

Velocities

5.1 Finite Element Formulation

We need to formulate a weak form of (4.5) so it can be solved using the finite element method. First we introduce a test function w, a partition of unity, into (4.1) giving

$$\int_{\Omega_c} w(h+\eta)d\Omega = C.$$

Then using a generalisation of the Reynolds Transport Theorem, and setting $\dot{\mathbf{x}} = \nabla p$ as before,

$$\int_{\Omega} w \nabla \cdot (h \nabla p) d\Omega = \int_{\Omega} w \nabla \cdot (h \mathbf{u}) d\Omega.$$

Integration by parts on both sides gives

$$-\oint_{\partial\Omega} wh\nabla p \cdot \mathbf{n} d\Gamma + \int_{\Omega} h\nabla w \cdot \nabla p d\Omega = -\oint_{\partial\Omega} wh\mathbf{u} \cdot \mathbf{n} d\Gamma + \int_{\Omega} h\nabla w \cdot \mathbf{u} d\Omega. \quad (5.1)$$

The boundary terms then cancel as they are either zero on the inner boundaries, or $\nabla p \cdot \mathbf{n} = \mathbf{u} \cdot \mathbf{n} = 0$ on the outer boundary. This condition

means that the mesh velocity has no component in the outward normal direction to the boundary, and so none of the mesh or the fluid leaves the boundary. We can also divide both sides by the constant h. The weak form (5.1) can then be expressed in finite element form by letting the continuous functions w and p be represented by the piecewise functions ϕ and P respectively, where ϕ is the P_1 basis function described in section 2.2. We then represent $P = \sum_j P_j \phi_j$. The finite element form is therefore

$$\int_{\Omega} \nabla \phi_i \cdot \nabla \phi_j d\Omega P_j = \int_{\Omega} \nabla w \cdot \mathbf{u} d\Omega,$$

which can be written in matrix vector form

$$K\mathbf{P} = \mathbf{f},\tag{5.2}$$

where K is a stiffness matrix and \mathbf{f} is a load vector.

5.1.1 Stiffness Matrix Entries

On the left hand side of (5.2), we have a stiffness matrix K. The entries of this matrix for each of the elements Ω_e are given by

$$K_{ij}^e = \int_{\Omega} \nabla \phi_i \cdot \nabla \phi_j d\Omega.$$

As $\nabla \phi_i$ and $\nabla \phi_j$ are constants, we can take these out of the integral. This then gives

$$K_{ij}^e = \nabla \phi_i \cdot \nabla \phi_j \int_{\Omega_e} d\Omega$$

where

$$\int\limits_{\Omega_e} d\Omega = \text{area of element} \equiv \Delta_{\text{area}}.$$

We also have that $\nabla \phi_i = |\nabla \phi_i| \mathbf{n}_i$, where \mathbf{n}_i is the normal vector in the direction of the node i.

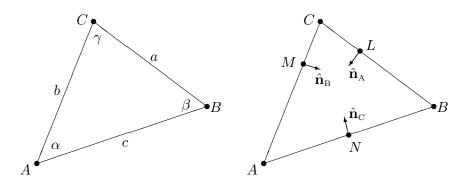


Figure 5.1: Dimensions of an element in the mesh

Looking at the element as in Figure 5.1, if we consider ϕ_B and ϕ_C , then

$$\nabla \phi_B \cdot \nabla \phi_C = |\nabla \phi_B| |\nabla \phi_C| \mathbf{n}_B \cdot \mathbf{n}_C$$
$$= \frac{1}{MB} \frac{1}{NC} cos(\pi - \alpha),$$

and the area of the element is

$$\Delta_{\text{area}} = \frac{aAL}{2} = \frac{ab\sin\gamma}{2}.$$

Therefore

$$K_{BC}^{e} = \frac{ab \sin \gamma \cos \alpha}{2ab \sin \gamma \sin \alpha}$$
$$= -\frac{\cot \alpha}{2}.$$

The other off diagonal entries in K are calculated similarly. The diagonal entries are found by noting that $\phi_A + \phi_B + \phi_C = 1$ everywhere, that is to say $\nabla \phi_A + \nabla \phi_B + \nabla \phi_C = 0$ everywhere. So

$$\int\limits_{\Omega_e} \nabla \phi_A \cdot \nabla \phi_A d\Omega \ = \ -\int\limits_{\Omega_e} \nabla \phi_A \cdot \nabla \phi_B d\Omega - \int\limits_{\Omega_e} \nabla \phi_A \cdot \nabla \phi_C d\Omega$$

$$=\frac{\cot\gamma}{2}+\frac{\cot\beta}{2}$$

Therefore we have the stiffness matrix for the element to be

$$K^{e} = \frac{1}{2} \begin{pmatrix} \cot \beta + \cot \gamma & -\cot \gamma & -\cot \beta \\ -\cot \gamma & \cot \alpha + \cot \gamma & -\cot \alpha \\ -\cot \beta & -\cot \alpha & \cot \alpha + \cot \beta \end{pmatrix}$$

Then this is assembled with the stiffness matrices for the whole domain, which are calculated in the same way.

5.1.2 Load Vector Entries

On the right hand side of the equation (5.2), we have a load vector \mathbf{f} . The entries of this vector for each of the elements are given by

$$f_i^e = \int_{\Omega_e} \nabla \phi_i \cdot \mathbf{u} d\Omega.$$

We can take the dot product of $\nabla \phi_i$ and \mathbf{u} outside the integral as this is a scalar constant for each element. We also know the flow velocities \mathbf{u} on the mid-segments of each element, so we can approximate $\mathbf{u} \approx \sum_{k=1}^{3} \mathbf{u}_k \psi_k$, as in Section 3.1. So for one element we have

$$f_i^e = \nabla \phi_i \cdot \left(\sum_{k=1}^3 \mathbf{u}_k \psi_k\right) \int_{\Omega_e} d\Omega,$$

where again $\int_{\Omega_e} d\Omega = \Delta_{\text{area}}$ and $\nabla \phi_i = |\nabla \phi_i| \hat{\mathbf{n}}_i$. If we again look at Figure 5.1 and consider ϕ_B , we have that

$$|\nabla \phi_B| = \frac{1}{BM} = \frac{1}{a \sin \gamma},$$

$$\hat{\mathbf{n}}_B = \frac{1}{BM} (x_A - x_L, y_A - y_L) = \frac{1}{a \sin \gamma} (x_A - x_L, y_A - y_L),$$

and

$$\Delta_{\text{\tiny area}} = \frac{ba\sin\gamma}{2}.$$

Therefore,

$$f_B^e = \frac{ba\sin\gamma}{2a^2\sin^2\gamma} (x_A - x_L, y_A - y_L) \cdot \left(\sum_{k=1}^3 \mathbf{u}_k \psi_k\right),$$

$$= \frac{b}{2a\sin\gamma} (x_A - x_L, y_A - y_L) \cdot \left(\sum_{k=1}^3 \mathbf{u}_k \psi_k\right).$$

The other entries in \mathbf{f} are calculated similarly and so we have

$$f_i^e = \Delta_{\text{area}} |\nabla \phi_i| \hat{\mathbf{n}}_i \cdot (\mathbf{u}_1 + \mathbf{u}_2 + \mathbf{u}_3),$$

as $\mathbf{u}_i \psi_i = \mathbf{u}_i$ for all *i*. This is then assembled over the whole domain with other load vector which are calculated similarly.

5.1.3 Mesh Velocities

Once the stiffness matrix and load vector have been computed, the system (5.2) can then be solved to find the mesh velocity potential, p, on each node using a conjugate gradient solver. Then we can find the mesh velocities. To get the mesh velocities, we need to find the gradient of the potential velocity at the nodes of the mesh as in (4.4).

This may be computed using a finite element formulation as follows. Multiplying by a test function w and integrating over Ω we get

$$\int_{\Omega} w \dot{\mathbf{x}} d\Omega = \int_{\Omega} w \nabla p d\Omega.$$

Then substituting the basis function ϕ_i for w, and approximating $p \approx P = \sum_j P_j \phi_j$ and $\dot{\mathbf{x}} \approx \dot{\mathbf{X}} = \sum_j \dot{\mathbf{X}}_j \phi_j$ gives

$$\sum_{j} \dot{\mathbf{X}}_{j} \int_{\Omega} \phi_{i} \phi_{j} d\Omega = \sum_{j} P_{j} \nabla \phi_{j} \int_{\Omega} \phi_{i} d\Omega, \tag{5.3}$$

which can then be solved for $\dot{\mathbf{X}}_j$ using a conjugate gradient solver.

However, we note that the integral on the left hand side of the equation gives the same mass matrix as in (3.7). To simplify the calculations this could again be lumped to give $\frac{\Delta_{\text{area}}}{3}I$ for each element, where I is the identity matrix. On the right hand side the integral is equal to $\frac{\Delta_{\text{area}}}{3}$.

Therefore we have that

$$\sum_{j} \dot{\mathbf{X}}_{j} = \sum_{j} P_{j} \nabla \phi_{j}.$$

To simplify this further, we consider the velocity computed on each element, which we will call the element velocity,

$$\dot{\mathbf{x}}_{\Omega_e} = \sum_{j=1}^3 P_j \nabla \phi_j.$$

However, this way of simplifying the calculation of the mesh velocity leads to a lot of averaging which may compromise the values of the mesh velocities. This is not as serious as it seems, since the main solver remains the Stommel equations.

Chapter 6

Implementation of the Mesh Velocities

We now discuss how the mesh velocities calculated in the previous chapter are implemented to move the mesh and the solution.

6.1 Interpolating the Mesh Velocities onto the P_1 nodes

The velocities have been found for each cell. We need to interpolate these onto the nodes of the mesh to enable nodal mesh movement. To do this we take an area weighted average of the element velocities in the elements surrounding each node (Figure 6.1) as follows:

$$\dot{\mathbf{x}}_i = \frac{\sum\limits_{e \in patch} \left| \Omega_e \right| \dot{\mathbf{x}}_{\Omega_e}}{\sum\limits_{e \in patch} \left| \Omega_e \right|}.$$

Now we have approximated the mesh velocities on the nodes we can use these to move the mesh.

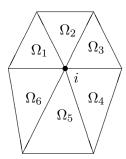


Figure 6.1: Averaging the element velocities to obtain the mesh velocities on the nodes

6.2 Moving the Mesh

We do not want to compute the mesh velocities every time step because of the small time step required for a stable solution due to the explicit time scheme. We will therefore move the mesh every n_{mesh} time steps, as follows:

$$\mathbf{x}^{n+1} = \mathbf{x}^n + \Delta t \, n_{\text{mesh}} \, \dot{\mathbf{x}}_{\text{mesh}}^n \tag{6.1}$$

We can do this because of the independence of the mesh equation from the solution equations.

6.3 Interpolating the Mesh Velocities onto the P_1^{NC} nodes

We now interpolate the mesh velocities onto the same P_1^{NC} nodes as the flow velocities on the mid-segments of each element. To do this we take a weighted average of the element velocities over two adjacent elements as follows:

$$\dot{\mathbf{x}}_i = \frac{\sum\limits_{e=1}^{2} |\Omega_e| \dot{\mathbf{x}}_{\Omega_e}}{\sum\limits_{e=1}^{2} |\Omega_e|}.$$

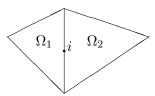


Figure 6.2: Averaging the element velocities to get the mesh velocities on the mid segments

These velocities can now be used to modify the solution.

6.4 Moving the Solution

When the mesh has been moved, we need to modify the solution so that it relates to the correct position in the mesh. After including the mesh adaptation procedure an extra term is required in each of the governing equations of the Stommel model as shown below.

By the chain rule for the conservation equation (2.3),

$$\left. \frac{\partial \eta}{\partial t} \right|_{m} = \left. \frac{\partial \eta}{\partial t} \right|_{s} + \dot{\mathbf{x}} \cdot \nabla \eta, \tag{6.2}$$

and similarly for the momentum equation (2.4),

$$\frac{\partial \mathbf{u}}{\partial t}\Big|_{m} = \frac{\partial \mathbf{u}}{\partial t}\Big|_{s} - \dot{\mathbf{x}} \cdot \nabla \mathbf{u}. \tag{6.3}$$

where m denotes the moving solution and s denotes the stationary solution.

The extra terms are advection terms, which counterbalance the movement of the mesh, and make sure that the change in mesh does not also change the solution.

Formulating the new terms using the finite element method with the same test functions as described in Chapter 3, gives, for the extra term in (6.2)

$$-\int_{\Omega} \phi_i \dot{\mathbf{x}} \cdot \nabla \eta d\Omega, \tag{6.4}$$

and for (6.3)

$$-\int_{\Omega} \phi_i \dot{\mathbf{x}} \cdot \nabla \mathbf{u} d\Omega. \tag{6.5}$$

We can then include (6.4) in f_{η} and (6.4) in $f_{\mathbf{u}}$, in equations (3.5) and (3.6) respectively, at the time step when the mesh is moved.

Chapter 7

Numerical Simulations and Results

As mentioned before, there are many methods for moving the mesh. We now look at several tests for moving the mesh. We set the mesh velocities:

- in a westward direction, given that we know the solution moves to the west.
- according to the conservation principle described in Chapter 4,
- equal to the flow velocities,
- a mixture of the two above tests.

For each case, we put these velocities into equation (6.1) to move the mesh, and also move the solution as described in Chapter 6.

7.1 Simple westward mesh movement

We start with a special example of setting the mesh velocities in the same direction as we know the solution is moving. We use the following mesh velocity equation

$$\dot{\mathbf{x}} = \left(\begin{array}{c} -\alpha x (1-x) \\ 0 \end{array} \right),$$

with different values of α . This is a simplified mesh movement, as we already know the direction that the solution will take. In other applications, we may not know how the solution evolves over time, and therefore would not be able to make this kind of assumption over which direction the mesh velocities should be travelling in.

7.1.1 Results for simple westward mesh movement

For all values of $\alpha > 0$ the mesh moves towards the west with varying speeds. When $\alpha = 0.5$, the mesh moves westwards very rapidly, even before the solution starts to move. This leads to instability in the mesh due to very narrow elements in the western region of the mesh and after 1693 time steps the solution blows up. Before this happens, however, the solution is not distorted by the mesh movement.

This leads us to look at smaller values of α to see if the solution remains stable as the mesh changes. Using $\alpha = 0.01$ gives better results in that, as well as there being no distortion in the solution, the solution remains stable up to maximum of 50,000 time steps that the solution is run for. This is not to say that it will always be stable for this value of α . The solution will eventually blow up as the elements in the western side of the mesh become very narrow, but takes longer to do

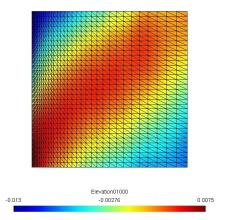


Figure 7.1: Solution after 1000 time steps with $\alpha = 0.5$.

so as $\alpha \to 0$.

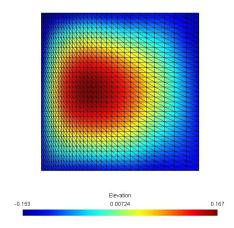


Figure 7.2: Solution after 50,000 time steps with $\alpha = 0.01$.

This simplified mesh movement is very easy to implement, but is a little crude. It is also not straightforward to move the mesh at a similar speed as the solution, and some sort of guess work is needed to determine the value of α to use to gain the best results. We know that the solution moves westward and a boundary layer forms on the

western boundary, but we do not know the speed at which the solution moves towards the west. Maybe a solution to these problems would be a variable α , which starts small as the solution starts rotating, then increases as the solution moves west, and finally decreases to zero as the boundary layer is formed and the westerly movement is stopped at the boundary.

7.2 Conservation based mesh movement

The second test is to use the mesh velocities that were described in Chapter 4. The principle here is one of conservation of the volume of water in a patch of elements over time. For this we have that

$$\dot{\mathbf{x}} = \nabla p.$$

In this case we expect the mesh to preserve the volume of water in a patch as the solution evolves through time. This means that patches of fluid over deeper regions may contain smaller elements than patches of fluid over shallower regions.

7.2.1 Results for conservation based mesh movement

The mesh velocity principle in Chapter 4 depends on the volume of water in each patch of elements being conserved over time. The effect of this conservation can be seen most at the start of the simulation. After 1000 time steps the elevation solution is a diagonally shaped band running from the south west to the north east of the region. This is due to the nature of the initial wind function; to the east in the north of the domain and to the west in the south of the domain. The

diagonal band has highest elevation in the centre of the diagonal, with decreasing elevation on either side.

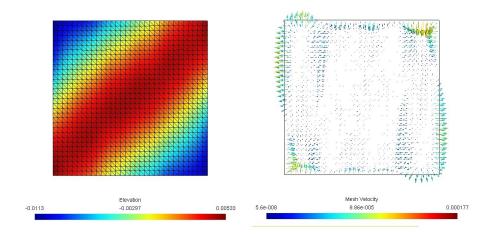


Figure 7.3: Solution after 1000 time steps using conservation principle

Note that in Figure 7.3 the velocities are not actually causing the mesh to leave the region. The length of the arrows has been exaggerated to show the direction of the mesh movement.

Using the conservation principle, we would expect patches of elements to be smaller in the area of high elevation, and larger in the area of low elevation. The results show mesh velocities directed into the high elevation in the north east and south west of the region. This suggests that the elements are being pushed together, and therefore reducing in size. In the low elevation areas, we see the mesh velocities moving away from these regions, suggesting that the elements are being pulled apart, or increasing in size. This is what we expected to happen when we used this principle.

As the elevation solution settles down into a rotating movement, the regions of high elevation are in the centre of the rotation, and the

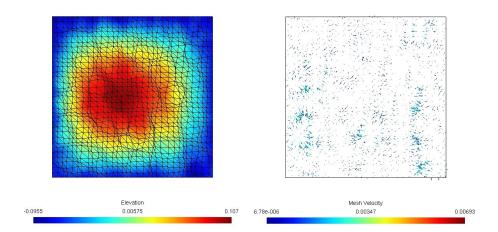


Figure 7.4: Solution after 25000 time steps using conservation principle

low elevation areas are near the boundaries. Because of the volume preservation, we would expect the patches of elements to be larger around the edges, and smaller in the centre of the rotation. Then as the solution moves westwards, we would expect the patches of elements in the western region to get smaller with the increased elevation, and patches of elements in the eastern region to grow. The results we get show that in the centre of the rotation, and around the boundaries, the patches are well preserved over time see Figure 7.4.

However, in the rest of the domain the conservation principle does not seem to have the same effect. This may be due to the number of nodes in the mesh being constant. This means that the patches that are growing in size are fighting for the space within the mesh, or the elements that are reducing in size are not doing so fast enough for the space to be created for the growing elements to fill. It could also be a result of the distortion in the mesh, leading to the solution becoming unstable, which may create more distortion in the mesh. The very narrow elements that are generated in parts of the mesh lead to the solution blowing up.

7.3 Flow based mesh movement

A purely Lagrangian method is to use the mesh velocity to be equal to the flow velocity. This would be the solution to the conservation principle described in Chapter 4 if we had imposed that $\nabla \times \dot{\mathbf{x}} = \nabla \times \dot{\mathbf{u}}$. We have that

$$\dot{\mathbf{x}} = \mathbf{u}$$
.

We expect the mesh to move round with the flow in a clockwise direction. This may not be ideal as the flow velocities on the boundary regions, particularly the western boundary, are much larger than the flow velocities on the central region. This may lead to an inappropriate tangled mesh.

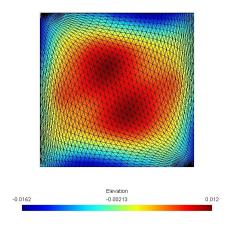


Figure 7.5: Solution after 3000 time steps using flow based mesh movement

7.3.1 Results for flow based mesh movement

The results confirm that the mesh moves in a clockwise direction around the centre of the region. The solution has been distorted by rotation as well. There is a smaller movement of nodes in the central region due to the smaller flow velocities here. The boundary condition $\mathbf{u} \cdot \mathbf{n} = 0$ means that the boundary nodes are stationary. This leads to very narrow elements in the four corners of the domain. The nodes become very close together on the boundary of the region, and the solution blows up after 3020 time steps.

7.4 Combined mesh movement

The last test we will look at is when the mesh velocities are calculated using a linear combination of the mesh velocities used in the previous two tests as follows,

$$\dot{\mathbf{x}} = \alpha \mathbf{u} + (1 - \alpha) \nabla p.$$

This corresponds to

$$\nabla \times \dot{\mathbf{x}} = \alpha \nabla \mathbf{u}.$$

We choose different values of α between 0 and 1. With a larger α we expect more rotation in the mesh than with a smaller value of α .

7.4.1 Results of combined mesh movement

Using $\alpha = 0.5$ gives results very similar to the results obtained when equating the mesh velocity to the flow velocity. This is because the flow velocities along the eastern and western boundary are larger in magnitude than the mesh velocities generated by the conservation principle,

whereas the flow velocity in the centre of the domain is comparable to the mesh velocities. We obtain the same narrow elements on the boundaries as before which cause the solution to blow up. Putting $\alpha=0.1$ gives slightly better results, but the flow velocities are still overpowering the conservation velocities, and again the solution blows up.

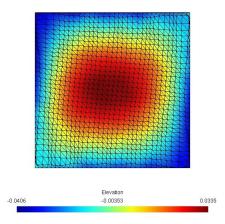


Figure 7.6: Solution after 9000 time steps using combined conservation/flow based mesh movement

By decreasing α the solution tends towards that of the conservation principle solution as we would expect. The solution also becomes more stable with the reduction in α .

Chapter 8

Conclusion

This dissertation has considered the possibility of incorporating a moving mesh method into an existing model of a special form of the shallow water equations, the Stommel model.

First we looked at how Stommel formed equations that describe the formation of a strong western boundary current such as the Gulf Stream. Then scale analysis was performed on these equations to determine the dominant terms. The result of this was that the rotational property of these velocities is an important part of the solution. This led us to investigate whether equating the mesh velocity to the flow velocities was a good strategy for moving the mesh. We saw in Chapter 7 that this resulted in a rotational mesh which became unstable very quickly. This is a result of the boundary nodes being restricted to remain on the boundary, as there is only flow velocity tangential to the boundary. Therefore the boundary nodes were able to move along the boundary which led to a bunching of nodes in the four corners of the domain.

We then discussed how the Stommel model was formulated using

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finite elements following [1]. These finite element equations were then discretised in time using an explicit Adams Bashforth method of order 3. The resulting mass matrix on the left hand side of the equation was lumped to enable us to avoid invertion of a matrix. Explicit methods, however, are not as stable as implicit methods, and a small time step of $\Delta t = 20$ seconds was required to keep a stable solution over the simulation. The initial conditions and mesh were then stated, followed by a description of how the solution evolves to form a western boundary layer over time. This formation of the boundary layer is one motivation for including a moving mesh procedure into the model.

In Chapter 4 a description of moving mesh methods was given, along with the benefits and drawbacks of using moving meshes. Then the use of a monitor function was introduced as in [8]. This dissertation has described the use of a conservation principle containing a monitor function together with a prescribed vorticity. This then gives rise to a conservation law from which we can obtain the mesh velocity. The conservation principle was applied to the chosen monitor function, fluid depth. In a finite element context, this prescribed that the volume of fluid in a patch of elements is constant for all time. The Reynolds Transport Theorem was then used along with Gauss' Theorem to change the equation from a Lagrangian to a Eulerian form in order to find the velocities. Then Helmholtz Decomposition Theorem was used and setting $\nabla \times \dot{\mathbf{x}} = 0$, to determine an equation for the mesh velocity $\dot{\mathbf{x}} = \nabla p$ in terms of the gradient of a velocity potential p.

The finite element formulation of this equation was described in Chapter 5. The calculation of the resulting stiffness matrix and load vector elements was described in terms of the dimensional properties of 44 Conclusion

each element. Then the equations were solved for the velocity potential using a conjugate gradient solver. A description of how $\dot{\mathbf{x}} = \nabla p$ could be solved using finite elements was then given. We then simplified the calculation by lumping the resulting mass matrix, so the use of a sparse matrix solver is not required. A further simplification was to find a velocity for each element which is constant over the whole element. This was then be interpolated onto the mesh nodes for mesh movement, and onto the mid-segment nodes for advecting the solution to compensate for the mesh movement. Whilst this simplification makes the computation easier, the process of calculating the mesh velocity for each element, and then having to interpolate this onto the mesh may have caused a slight loss of accuracy in the mesh velocities, but should not alter the solution.

The interpolation equations for generating nodal, and mid-segment velocities from the element velocities was described in Chapter 6. Averages were weighted according to the areas of the elements to try to obtain more accurate interpolations to the node velocities and mid-segment velocities. This averaging is not ideal as we could have calculated the mesh velocity straight from the finite element form in Section 5.2, and then used the average of the end points of a segment to determine the mid-segment velocities. However, use of computational time was an issue because of the small time step required for a stable solution. Therefore we chose the simplifications over the perhaps slight increase in accuracy. The equations of mesh movement were then stated. The mesh velocity was not calculated at every time step as the reduced time step means that the solution did not move very quickly over the simulation. The extra terms that arise in the solution

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equations at the time step when the mesh movement occurs are stated, followed by the finite element formulation of these terms.

Chapter 7 discussed the results of numerical simulations for a number of test cases for different mesh movement ideas. These ideas were based on the ideas generated throughout the dissertation. The first was a special case, as the Stommel model has an analytic solution so we know that the elevation field flows westwards, considering purely westward movement of the mesh. This gave a very simple computation of the mesh velocity, but was also very crude, and the solution blew up because of the gathering of very narrow elements on the western boundary. A modified version of this idea is discussed in Chapter 9. The main idea in this dissertation was to conserve the depth function over a patch of elements with a prescribed mesh velocity. The results show that this is been a promising start to describing a moving mesh procedure for these equations. The time restriction for this dissertation means that other monitor functions or curl properties of the depth monitor function have not been able to be investigated. Some ideas for other curl properties or monitor functions are in Chapter 9. Another test was to set the mesh velocity to be equal to the flow velocity, equivalent to choosing the vorticity of the mesh to be equal to the vorticity of the fluid. This resulted in a very twisted mesh and an unstable solution. The final idea was to take a linear combination of the conservation based velocities and the flow velocity. The results showed that the flow velocity overpowered the conservation velocities, and a very small ratio of flow to conservation based mesh movement was required for the solution to remain stable for any length of time.

Chapter 9

Further Work

The scope of work in this dissertation is limited by time available to carry out the work. Given more time, there are other areas that would be interesting to investigate. Several of these are given below.

More complex equations for western mesh movement Perhaps an improved equation to use to move the mesh west would include a term that would allow for the size of the boundary layer, using the analytic solution to the Stommel equations. This could be used to make sure the size of the elements in this boundary layer do not become too small. This would of course only apply for the Stommel problem and not for general equations.

Other monitor functions The monitor function described in this dissertation is just one of many that could be used as a basis for mesh movement. Use of different monitors that may give better results could be investigated. Other monitor functions that could be used include

• vorticity or potential vorticity

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• fluid surface gradient,

$$|\nabla \eta|$$

• area on the surface of the solution,

$$\sqrt{1+(\nabla\eta)^2}$$

The latter two would be more suitable for moving points in areas with large gradients, as in a boundary layer

Differing curl properties of mesh velocity We could investigate other curl properties using $\nabla \times \dot{\mathbf{x}} = \nabla \times q$, where q is some variable used to determine the mesh velocity, along with a monitor function. In this case, $\dot{\mathbf{x}}$ is not only the gradient of a potential function, but

$$\dot{\mathbf{x}} = q + \nabla p.$$

For the monitor function we have used, we have looked at the cases when; $\nabla \times \dot{\mathbf{x}} = 0$ taking q = 0, $\nabla \times \dot{\mathbf{x}} = \nabla \times \mathbf{u}$ where $q = \mathbf{u}$, and $\nabla \times \dot{\mathbf{x}} = \alpha \mathbf{u}$ which is equivalent to taking $q = \alpha \mathbf{u}$.

Modelling tides with a moving mesh. The possibility of using a moving mesh to show the rising and falling of the coastlines with tides could also be looked at. This could be done by prescribing the surface elevation on the boundaries to rise and fall over time. The boundary condition would need to be changed from $\mathbf{u} \cdot \mathbf{n} = 0$ on $\partial \Omega$ to $\mathbf{u} \cdot \mathbf{n} = \dot{\mathbf{x}} \cdot \mathbf{n}$ to enable the mesh to move in and out of the domain.

Grid generation Different initial grids could be generated using an equidistribution procedure. As the initial solution for η and \mathbf{u} is zero

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over the whole domain, the grid could be generated using the initial wind stress. An alternative would be to include a feature such a Gaussian hill into the domain so that the initial elevation field is not zero, and the mesh could be generated by equidistributing the volume of fluid over the domain.

Other applications The principles outlined in this dissertation could be used to introduce mesh movement into other models of fluid flow, such as the Navier Stokes equations.

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